finally calcareous siliciclastics. This same pattern occurs cyclically as a vertical sequence with oil shale at the base.

The calcimicrites at the depocenter increase in dolomite content dramatically in the upper third of the section, and oil shale units become more widespread. There is an accompanying increase in saline minerals as well as zeolites indicating hypersalinity.

Bioturbation is absent from both oil shale and dolomicrites. The marginal bioturbated facies are always dominated by calcimicrite, even in the upper third of the section. Well-preserved fossil fish occur in kerogen-rich calcimicrites, but never in dolomicrites.

The facies patterns and vertical sequences in Fossil basin are interpreted as being deposited in a closed-basin occupied by a lake with an ephemeral and hypersaline hypolimnion that underwent frequent vertical fluctuations that affected large areas of the lake bottom. The "transgression-regression" of the hypolimnion over large areas of lake bottom was possible because of a low topographic gradient on the lake bottom. Deposition of oil shale occurred within the denser, cooler, and hypersaline hypolimnion waters—laminated calcimicrite in a zone of chemocline fluctuation—and bioturbated calcimicrite was deposited in fresher nearshore zones. Kerogen-rich carbonates became more widespread in the later stages of the lake when it evolved into a shallow hypersaline lake. This suggests that kerogen deposition was controlled by high salinity rather than anoxic conditions.

It is concluded that the carbonate facies patterns, vertical sequences, and oil shale genesis required a dynamic and fluctuating lake as well as a fluctuating chemocline level in a closed basin with a low topographic gradient. This differs from the static, anoxic, and deep basin model frequently postulated for Green River Formation oil shale deposition.

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Recurrent Motion on Precambrian-Age Basement Faults, Palo Duro Basin, Texas Panhandle

The distribution of Late Precambrian(?) through Quaternary strata in the Palo Duro basin and surrounding uplifts documents recurrent motion on Precambrian-age basement faults. Basement blocks have been uplifted with little tilting or folding of overlying strata along a system of northwest-southeast oriented faults, part of a regional trend extending from central Colorado to southwestern Oklahoma. The orientation of basement terranes in Colorado and that of a 50-mi (80-km) long mylonite zone in east-central New Mexico suggest a Precambrian age for the faults.

An Arkosic sandstone overlies basement and underlies a Cambrian(?) quartzose sandstone in a few Palo Duro basin wells. It may represent debris shed from active fault blocks during the opening of the southern Oklahoma aulocogen in the Late Precambrian or Early Cambrian. Ordovician carbonates thin or are missing beneath Mississippian carbonates on some fault blocks, indicating a post-Ordovician-pre-Mississippian period of faulting

The greatest amount of deformation occurred during the Pennsylvanian. Thickness, distribution, and facies of sediments were controlled by the location of active faults. Lower Pennsylvanian strata thin by up to 50% across some structures. Fault blocks provided sources of arkosic debris and loci for carbonate buildups throughout the Pennsylvanian and Early Permian. Around the periphery of the basin, Late Pennsylvanian or Early Permian faulting caused a wedging out of older units beneath the Wolfcamp.

Permian, Triassic, and Neogene units, along with present topography, all have been subtly affected by basement structures. The entire section thins over basement highs. Middle and Upper Permian evaporites are thicker in structural lows. The overlying Dockum Group (Triassic) and Ogallala Formation (Neogene), both nonmarine clastic units, become finer grained over basement highs. Present topographic highs coincide with some basement highs. Also, in some places remarkably straight stream segments parallel basement faults. Low-level seismic activity, primarily north and west of the Palo Duro basin, suggests continuing motion on at least some of the faults.

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Results of Ocean Margin Drilling Program Synthesis of Gulf of Mexico Basin

A series of 23 maps plus three cross sections synthesize and integrate for the first time geologic and geophysical data from both the deep central Gulf of Mexico basin and the periphery of the basin. These maps and sections are part of the Ocean Margin Drilling Program synthesis of the Gulf of Mexico basin, a joint project sponsored by the National Science Foundation/Joint Oceanographic Institutions, Inc., and a consortium of petroleum companies. The study area is bounded by 30°/31° on the north, 98° on the west, 82° on the east, and 18° on the south (excluding Cuba and the Yucatan basin). Maps include a regional tectonic map; a map of all seismic refraction data; six structure maps (basement, top Jurassic, top Early Cretaceous, top Late Cretaceous, top Paleogene, and top Neogene); six lithofacies maps (Oxfordian [Late Jurassic], Aptian-Cenomanian [Early Cretaceous], Coniacian-Santonian [Lake Cretaceous], Early Eocene, Miocene, and Pleistocene); and nine isopach maps (total sediment, pre-top Early Cretaceous, post-top Early Cretaceous, Late Jurassic, Early Cretaceous, Lake Cretaceous, Paleogene, Neogene, and Pleistocene). Also included are two N-S cross sections and one E-W cross section. This project is an attempt to synthesize all data available in the public domain. Data for the deep central Gulf are based mainly on regional multifold seismic lines, while data from the periphery are based mainly on the published literature. These maps and sections present for the first time an integrated and comprehensive look at the structure and stratigraphy of the entire central Gulf basin as well as document many details of the early history and later filling of the basin.

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Dolomite Selectivity, An Experimental Approach

Hydrothermal dolomitization of corals, gastropods, pelecypods, echinoderms, forams, and coralline algae indicates that grain size is more important than mineralogy in determining (1) whether or not a fossil will be dolomitized and (2) whether or not the dolomite will pseudomorphically replace a fossil.

Dolomite commonly selectively replaces matrix and/or specific fossils. When dolomite replaces fossils, certain fossils retain their optical characteristics (i.e., pseudomorphic replacement). These selective characteristics have been attributed to both grain size and mineralogy and have been used to make inferences about the predolomitization diagenetic history of sediments.

Fossils were dolomitized at 250°C (482°F) in Ca/MgCl solutions for periods of time from 4.5 to 320 hours. Aragonitic corals, gastropods, and pelecypods formed stoichiometric, microcrystalline, xenotopic dolomite and low Mg-calcite (LMC). The dolomite was not pseudomorphic after the aragonite. The conversion of AR \rightarrow LMC is more rapid than the AR \rightarrow DOL in these experiments. For instance, gastropods run for 23 hours formed dolomite and LMC in a ratio of 1/10, at 170 hours the ratio was 1/4, and at 340 hours the ratio was 1/1.

HMC coralline algae, forams, and echinoderm fragments were dolomitized before and after conversion to LMC. The dolomite formed was cryptocrystalline and pseudomorphic after the forams and echinoderms regardless of the mineralogy. We attributed this to the cryptocrystalline nature of the substrate.

Oyster fragments composed of microcrystalline LMC formed non-stoichiometric, poorly ordered dolomite even after 320 hours. None of the other reactants were as resistant to dolomitization.

Our results indicate that grain size is more important than mineralogy in determining the fabric of dolomite replacement crystals. Both HMC and LMC can be pseudomorphically replaced. Pseudormophic replacement requires (1) abundant nucleation sites and (2) a regular crystallographic relationship between the calcite and dolomite. Argonite was not pseudomorphically replaced, probably because it was microcrystalline rather than cryptocrystalline. Also, most of the aragonite converted to LMC prior to dolomitization.

Selective replacement characteristics of many natural dolomites are readily explained as being an effect of the grain size of the material replaced. Freshwater diagenesis of a sediment prior to dolomitization may retard dolomitization if the grain size of the CaCO₃ is increased. However, conversion of LMC without appreciable increase in grain size may not retard dolomitization.

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Petroleum Exploration and Resource Potential of Offshore Newfoundland and Labrador

The continental margin of Newfoundland and Labrador, encompassing a total area of 714, 000 mi² (1,849,252 km²) has been the target of exploratory activity since the early 1960s.

Exploratory drilling began on the Grand Banks in 1966 and by 1974 a total of 40 dry wells had been drilled. This lack of success, accompanied by escalating drilling costs, resulted in the curtailment of exploratory activities. In 1979 wildcat drilling resumed on the Grand Banks and the Hibernia field was discovered with the drilling of the P-15 well. This well, with an estimated flow potential of more than 20,000 BOPD, was the first oil well drilled on the Atlantic shelf of North America capable of commercial production. Truly a "giant," the Hibernia structure has a resource potential of 1.85 billion bbl of oil and 2.0 tcf of gas at a probability level of 50%. Six significant oil discoveries have been made on the Grand Banks. Of these, the Hibernia, Nautilus, Hebron, and Ben Nevis discoveries are located in highly faulted hinge zones on the western and eastern flanks of the northward plunging Avalon basin graben. The South Tempest structure is located on a ridge complex to the east of the Avalon basin. The Adolphus well drilled a salt piercement structure in the basin depocenter. The reservoirs are fluvial-deltaic and shoreline sandstones of Jurassic and Cretaceous age.

Since 1971, 25 wells drilled on the Labrador Shelf resulted in one oil and five gas discoveries. The reservoirs are Paleozoic carbonates and Lower Cretaceous, Paleocene, and Eocene sandstones. All are capping or draping basement horst blocks.

By the end of 1982, total exploratory efforts had resulted in the drilling of 86 wells and the acquisition of approximately 240,000 line-mi of marine reflection seismic. Provincial land permits on the continental margin are held by ten permittees. This land position represents 54 million acres (22 million ha.) and 133 exploratory permits. A total resource potential of these structures has been estimated at 14.7 billion bbl of oil and 88.6 tcf of gas at a 50% probability level. A commercial discovery was long in coming but the recent high success rates confirm this margin as a major frontier of enormous potential.

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Regional Distribution of Hydrocarbon Fluid Inclusions in Carbonate Fracture Filling Cements: Geohistory Analysis and Timing of Oil Migration, Oman Foredeep

Fractured, reservoir limestones in Oman and the United Arab Emirates include the Shuaiba (lower Aptian) and Mauddud (upper Aptian-lower Cenomanian). Deposition of these bioturbated, argillaceous foraminiferal-peloidal wackestones and packstones ceased in the Early Cretaceous as the Oman foredeep subsided and filled with pelagic sediment. Petrography and geohistory analysis of four wells and one outcrop suite reveals five stages of diagenesis, fracturing, and fluid migration. (1) Shelf emergence: early cementation associated with regional unconformities overlying both limestones; (2) pre-orogenic shelf emergence, late Cenomanian to Turonian: fractures cutting Stage 1 cements are healed by very cloudy, cleaved, and twinned calcite containing microfractures with yellow-white fluorescent, hydrocarbon fluid inclusions; (3) initial foredeep downwarp of 0 to 800 m (0 to 2,624 ft), Coniacian to early Campanian: fractures crosscutting Stage 2 fractures are healed with cloudy, cleaved, and sometimes twinned calcite containing dull-blue fluorescent, hydrocarbon fluid inclusions; (4) rapid subsidence and filling with 600 to 3,400 m (1,970 to 11,155 ft) of flysch, exotic blocks, and thrust toes, Campanian to Maestrichtian: burial and tectonic stylolites crosscut Stage 2 and 3 fractures; and (5) uplift of the Oman Mountains after 3,900 + m (12,795 + ft) burial by early Tertiary: fractures crosscutting all diagenetic features are filled with clear untwinned and uncleaved calcite containing only nonfluorescent, aqueous fluid inclusions. If we can correlate earliest stylolite formation with a minimum burial load of ~800 m $(\sim 2,625 \text{ ft})$, then the hydrocarbon inclusions in Stage 2 fractures must predate all of Stage 4 and most of Stage 3. In the deepest portions of the foredeep, close to the Oman Mountain front, this limits the presence of oil in fracture porosity to late Turonianearly Campanian time. Farther to the west, in the shallower parts of the foredeep, this constraint relaxes, and oil migration occurred as late as early Tertiary.

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Illite/Smectite Diagenesis and Hydrocarbon Generation in Cretaceous Mowry and Skull Creek Shales of Northern Rocky Mountains-Great Plains Region

The Lower Cretaceous Mowry and Skull Creek Shales and their equivalents are among the major source rocks in the northern Rocky Mountains-Great Plains region. They are the major